

# Data Assimilation in Atmospheric CTMs: I. Computational Tools

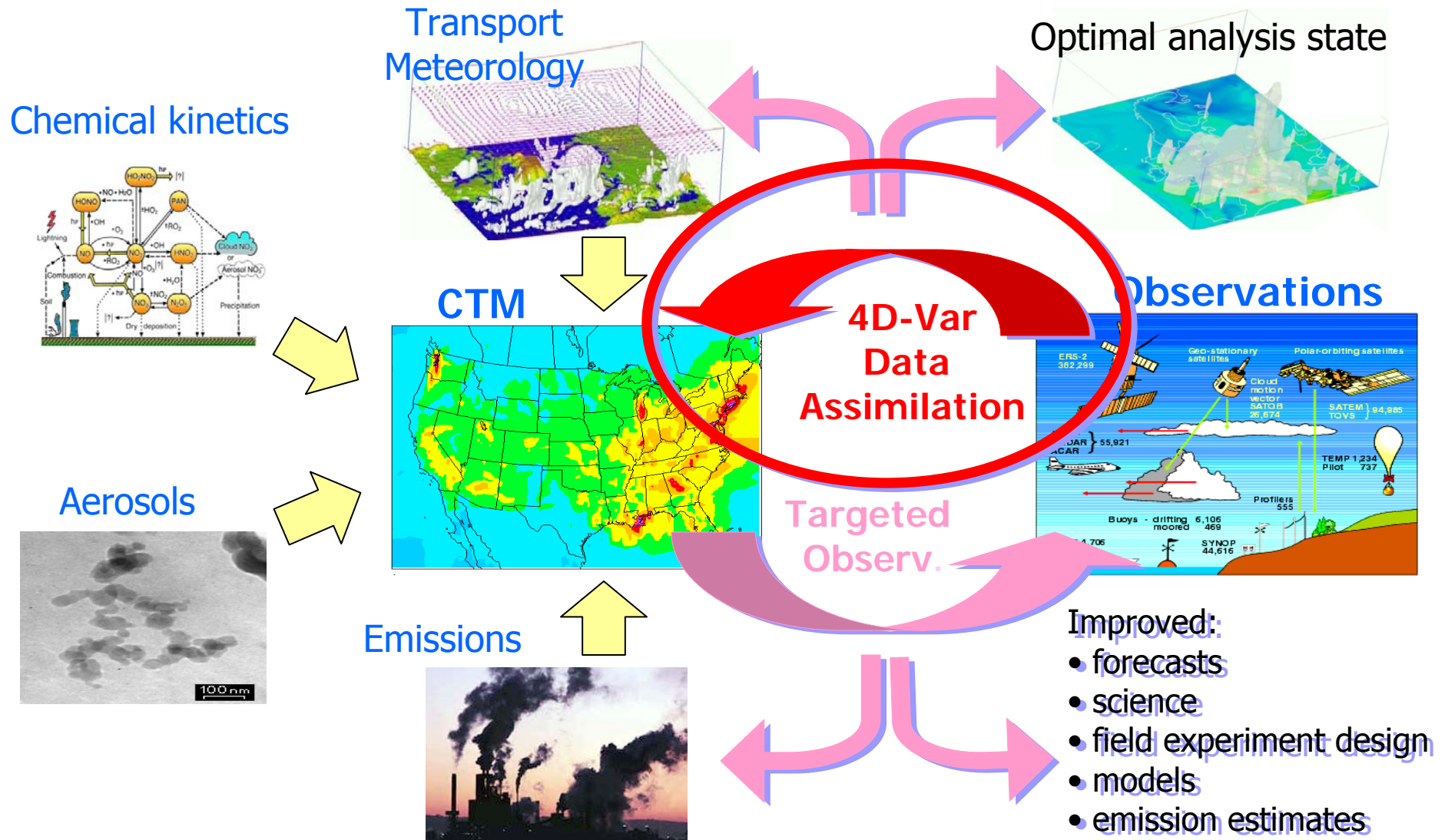
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(Virginia Tech)

G.R. Carmichael et al. (U. Iowa)

J.H. Seinfeld et al. (Caltech)

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# Information feedback loops between CTMs and observations: data assimilation and targeted meas.



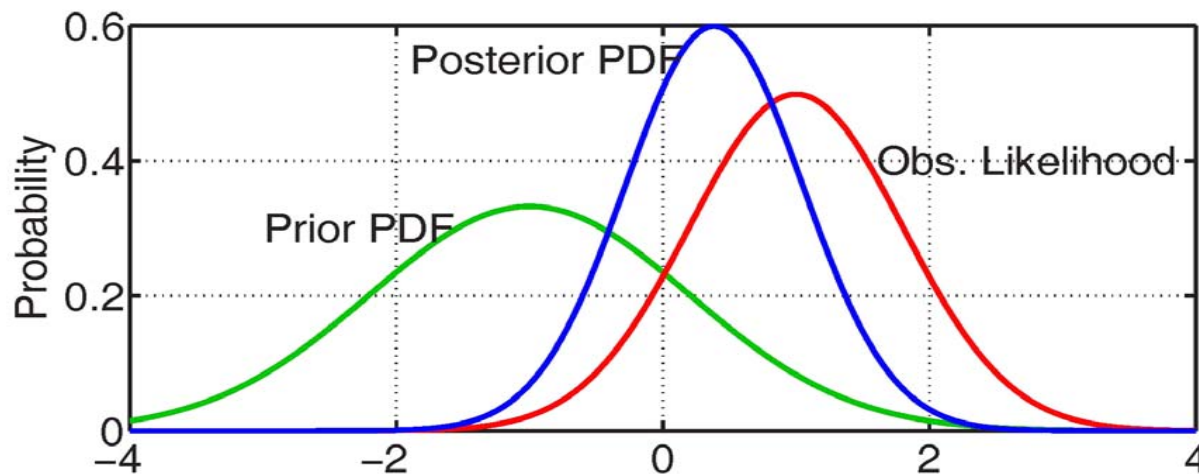
# Best estimate of the state/parameters is obtained by combining multiple sources of information

**Model** (encapsulating knowledge on the physics, chemistry, thermodynamics, etc)

**Background** (encapsulating best a-priori knowledge of the state)

**Observations** (encapsulating new information about reality)

**Bayes:** 
$$P\left[y^k \mid z_{\text{obs}}^k \dots z_{\text{obs}}^0\right] = \frac{P[z_{\text{obs}}^k \mid y^k] \cdot P[y^k \mid z_{\text{obs}}^{k-1} \dots z_{\text{obs}}^0]}{\int P[z_{\text{obs}}^k \mid y] \cdot P[y \mid z_{\text{obs}}^{k-1} \dots z_{\text{obs}}^0] dy} = \frac{P_R[\varepsilon_{\text{obs}}^k] \cdot P_B[y^k]}{\dots}$$



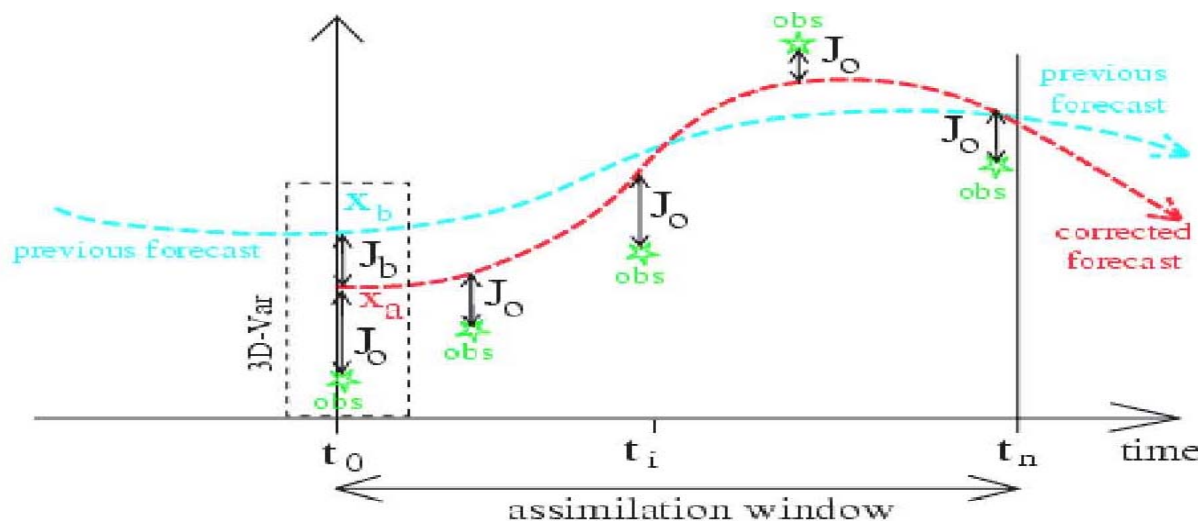
**Methods:**  
4D-Var, EnKF

*[Picture from  
J.L. Anderson]*

In the 4D-Var approach D.A. is formulated as a PDE-constrained optimization problem (gradient-based)

$$\min_{\mathbf{y}^0} \psi(\mathbf{y}^0) = \frac{1}{2} (\mathbf{y}^0 - \mathbf{y}^b)^T \mathbf{B}^{-1} (\mathbf{y}^0 - \mathbf{y}^b) + \frac{1}{2} \sum_{k=1}^N (\mathbf{H}^k \mathbf{y}^k - \mathbf{z}^k)^T \mathbf{R}_k^{-1} (\mathbf{H}^k \mathbf{y}^k - \mathbf{z}^k)$$

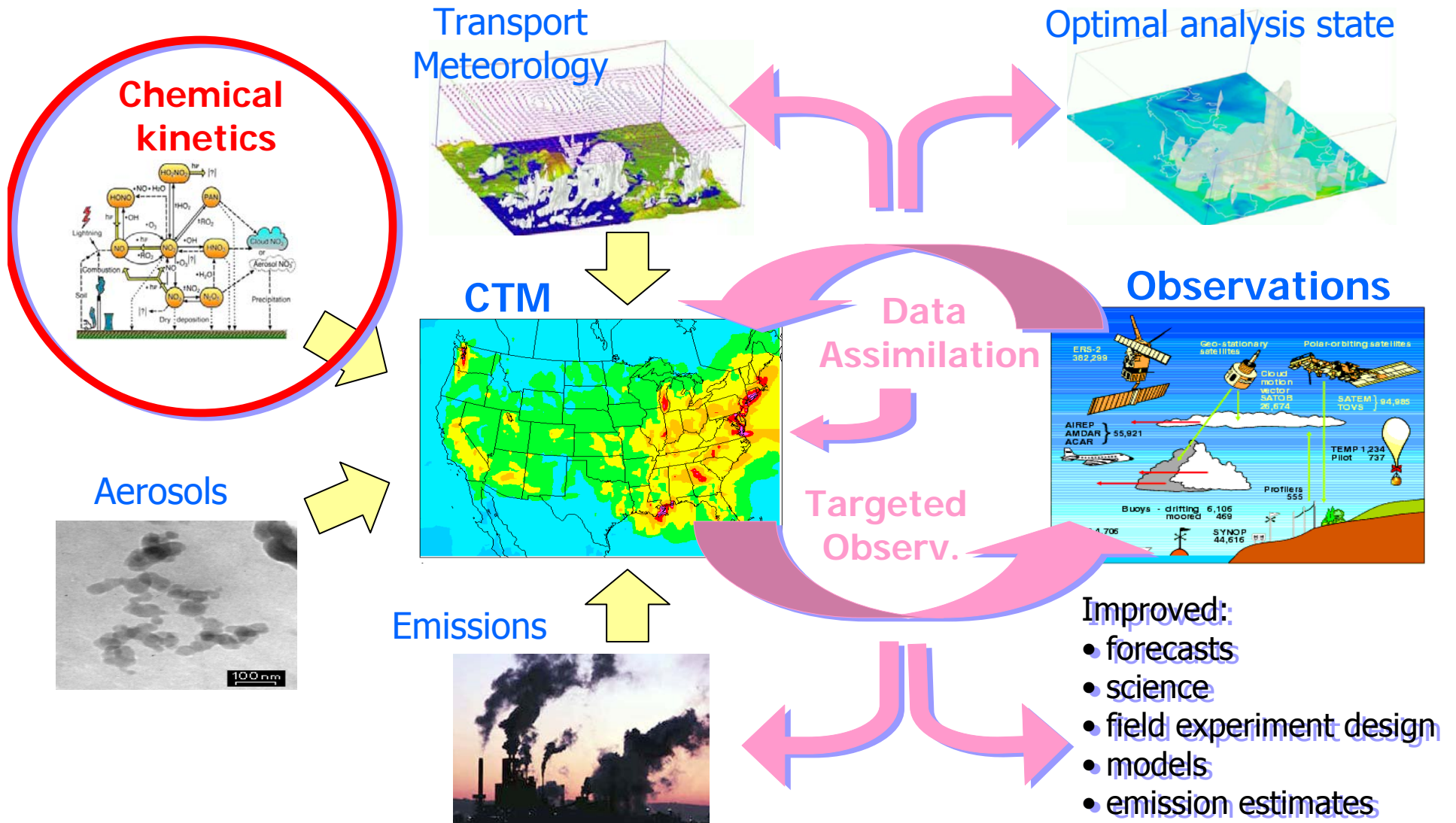
subject to  $\mathbf{y}^k = M(t^{k-1}, \mathbf{y}^{k-1}, \mathbf{p}), \quad k = 1, 2, \dots$



[Picture from  
F. Bouttier, P. Courtier]

**Gradient:**  $\lambda = \nabla_{\mathbf{y}^0} \psi = \left( \frac{\partial \psi}{\partial \mathbf{y}^0} \right)^T = \mathbf{B}^{-1} (\mathbf{y}^0 - \mathbf{y}^b) + \sum_{k=1}^N \left( \frac{\partial \mathbf{y}^k}{\partial \mathbf{y}^0} \right)^T (\mathbf{H}^k)^T \mathbf{R}_k^{-1} (\mathbf{H}^k \mathbf{y}^k - \mathbf{z}^k)$

# Adjoins of stiff chemical kinetics: formulation, challenges, and automatic implementation



# KPP automatically generates simulation and direct/adjoint sensitivity code for chemistry

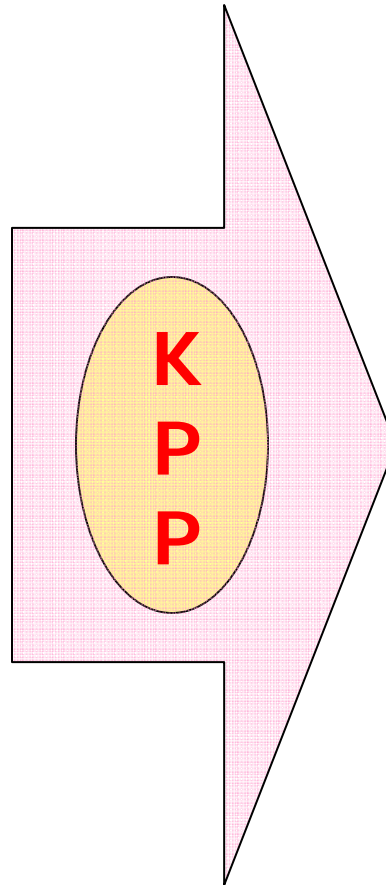
## Chemical mechanism

```
#INCLUDE atoms

#DEFVAR
O = O; O1D = O;
O3 = O + O + O;
NO = N + O;
NO2 = N + O + O;

#DEFFIX
O2 = O + O; M = ignore;

#EQUATIONS { Small Stratospheric }
O2 + hv = 2O      : 2.6E-10*S;
O  + O2 = O3      : 8.0E-17;
O3 + hv = O  + O2 : 6.1E-04*S;
O  + O3 = 2O2     : 1.5E-15;
O3 + hv = O1D + O2 : 1.0E-03*S;
O1D + M = O  + M  : 7.1E-11;
O1D + O3 = 2O2     : 1.2E-10;
NO  + O3 = NO2 + O2 : 6.0E-15;
NO2 + O  = NO  + O2 : 1.0E-11;
NO2 + hv = NO  + O  : 1.2E-02*S;
```



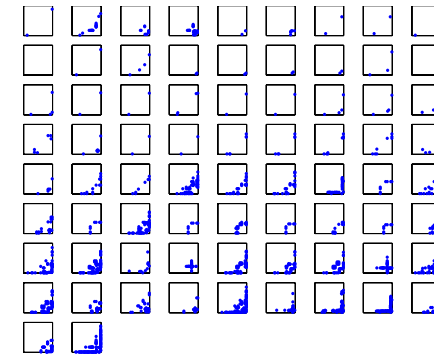
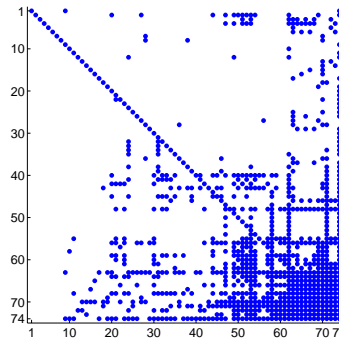
## Simulation code

```
SUBROUTINE FunVar ( V, F, RCT, DV )
  INCLUDE 'small.h'
  REAL*8 V(NVAR), F(NFIX)
  REAL*8 RCT(NREACT), DV(NVAR)
  C A - rate for each equation
  REAL*8 A(NREACT)
  C Computation of equation rates
  A(1) = RCT(1)*F(2)
  A(2) = RCT(2)*V(2)*F(2)
  A(3) = RCT(3)*V(3)
  A(4) = RCT(4)*V(2)*V(3)
  A(5) = RCT(5)*V(3)
  A(6) = RCT(6)*V(1)*F(1)
  A(7) = RCT(7)*V(1)*V(3)
  A(8) = RCT(8)*V(3)*V(4)
  A(9) = RCT(9)*V(2)*V(5)
  A(10) = RCT(10)*V(5)
  C Aggregate function
  DV(1) = A(5)-A(6)-A(7)
  DV(2) = 2*A(1)-A(2)+A(3)-A(4)+A(6)-&A(9)+A(10)
  DV(3) = A(2)-A(3)-A(4)-A(5)-A(7)-A(8)
  DV(4) = -A(8)+A(9)+A(10)
  DV(5) = A(8)-A(9)-A(10)
  END
```

[Damian et.al., 1996; Sandu et.al., 2002]

# Rosenbrock, Runge-Kutta, Sdirk methods and their adjoints are efficiently implemented by KPP

SAPRC-99  
Sparse Jacobian  
and Hessian



Rosenbrock  
Method  
( $T_{\text{fwd}}$ )

$$\begin{cases} \mathbf{y}^{n+1} = \mathbf{y}^n + \sum_{j=1}^s m_j \mathbf{k}_j, & \mathbf{Y}^i = \mathbf{y}^n + \sum_{j=1}^{i-1} a_{i,j} \mathbf{k}_j \\ \left[ \frac{1}{h\gamma} \mathbf{I} - \mathbf{J}^n \right] \cdot \mathbf{k}_i = \mathbf{f}(\mathbf{Y}^i) + \sum_{j=1}^{i-1} \frac{1}{h} c_{i,j} \mathbf{k}_j, & 1 \leq i \leq s \end{cases}$$

Discrete  
Adjoint  
( $T \approx 2.3T_{\text{fwd}}$ )

$$\begin{cases} \left[ \frac{1}{h\gamma} \mathbf{I} - (\mathbf{J}^n)^T \right] \cdot \mathbf{u}_i = m_i \lambda^{n+1} + \sum_{j=i+1}^s \left( a_{j,i} \mathbf{v}_j + \frac{1}{h} c_{j,i} \mathbf{u}_j \right), & \mathbf{v}_i = \mathbf{J}^T(\mathbf{Y}^i) \cdot \mathbf{u}_i \\ \lambda^n = \lambda^{n+1} + \sum_{i=1}^s \left( \mathbf{H}^n \times \mathbf{k}_i \right)^T \cdot \mathbf{u}_i + \sum_{i=1}^s \mathbf{v}_i \end{cases}$$

[Sandu et.al., 2002]

# Runge-Kutta methods and their adjoints are well suited for inverse chemical kinetic problems

RK Method 
$$\mathbf{y}^{n+1} = \mathbf{y}^n + h \sum_{i=1}^s b_i \mathbf{f}(\mathbf{Y}^i), \quad \mathbf{Y}^i = \mathbf{y}^n + h \sum_{j=1}^s a_{i,j} \mathbf{f}(\mathbf{Y}^j)$$

Continuous Adjoint 
$$\lambda^n = \lambda^{n+1} + h \sum_{i=1}^s b_i \mathbf{J}^T(\mathbf{y}^{n+1-c_i h}) \cdot \Lambda^i, \quad \Lambda^i = \lambda^{n+1} + h \sum_{j=1}^s a_{i,j} \mathbf{J}^T(\mathbf{y}^{n+1-c_i h}) \cdot \Lambda^j$$

Discrete Adjoint 
$$\lambda^n = \lambda^{n+1} + \sum_{i=1}^s \theta^i, \quad \theta^i = h \mathbf{J}^T(\mathbf{Y}^i) \cdot \left[ b_i \lambda^{n+1} + \sum_{j=1}^s a_{j,i} \theta^j \right]$$
  
*[Hager, 2000]*

**Consistency:** The discrete adjoint of RK method of order  $p$  is an order  $p$  discretization of the adjoint equation. (Proof using elementary differentials of transfer functions).

**Stiff behavior.** For SPP apply RK with invertible coefficient matrix  $A$  and  $R(\infty) = 0$ . If the cost function depends only on the non-stiff variable  $y$  then  $\lambda_z = 0$  and  $\lambda_y$  are solved with the same accuracy as the original method, within  $\mathcal{O}(\epsilon)$ .

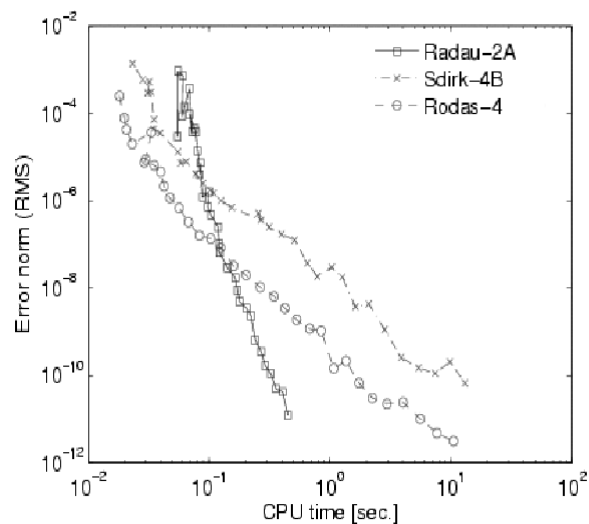
*[Sandu et al., 2005]*



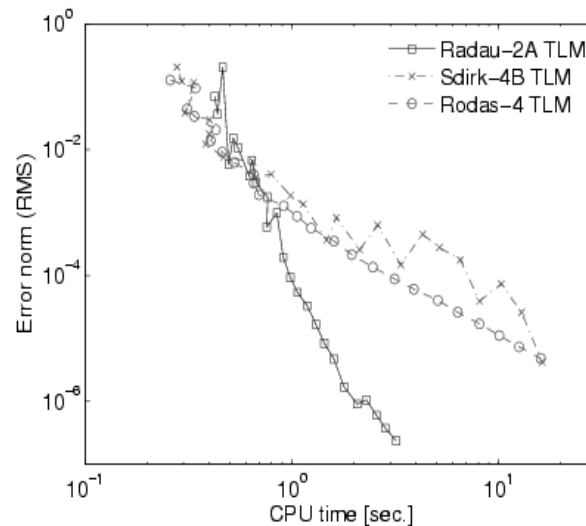
# Methods available in the KPP numerical library

- **FIRK** 3-stage: Radau-2A (ord.5), Radau-1A (ord.5), Lobatto-3C (ord.4), Gauss (ord.6)
- **SDIRK**: 2a, 2b (2s, ord.2), 3a (3s, ord.2), 4a, 4b (5s, ord.4)
- **Rosenbrock**: Ros2, Ros3, Ros4, Rodas3, Rodas4.

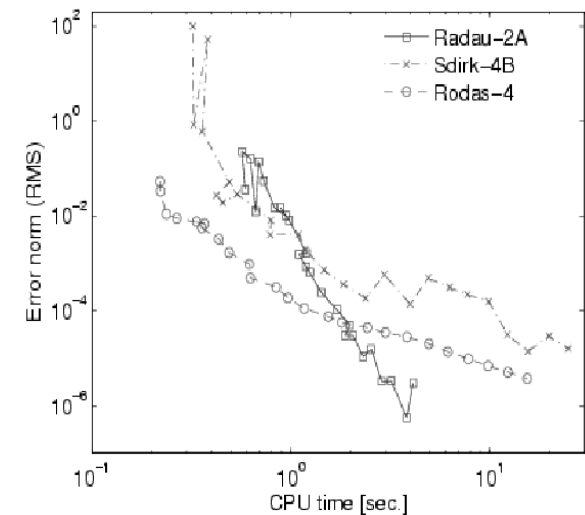
## Forward



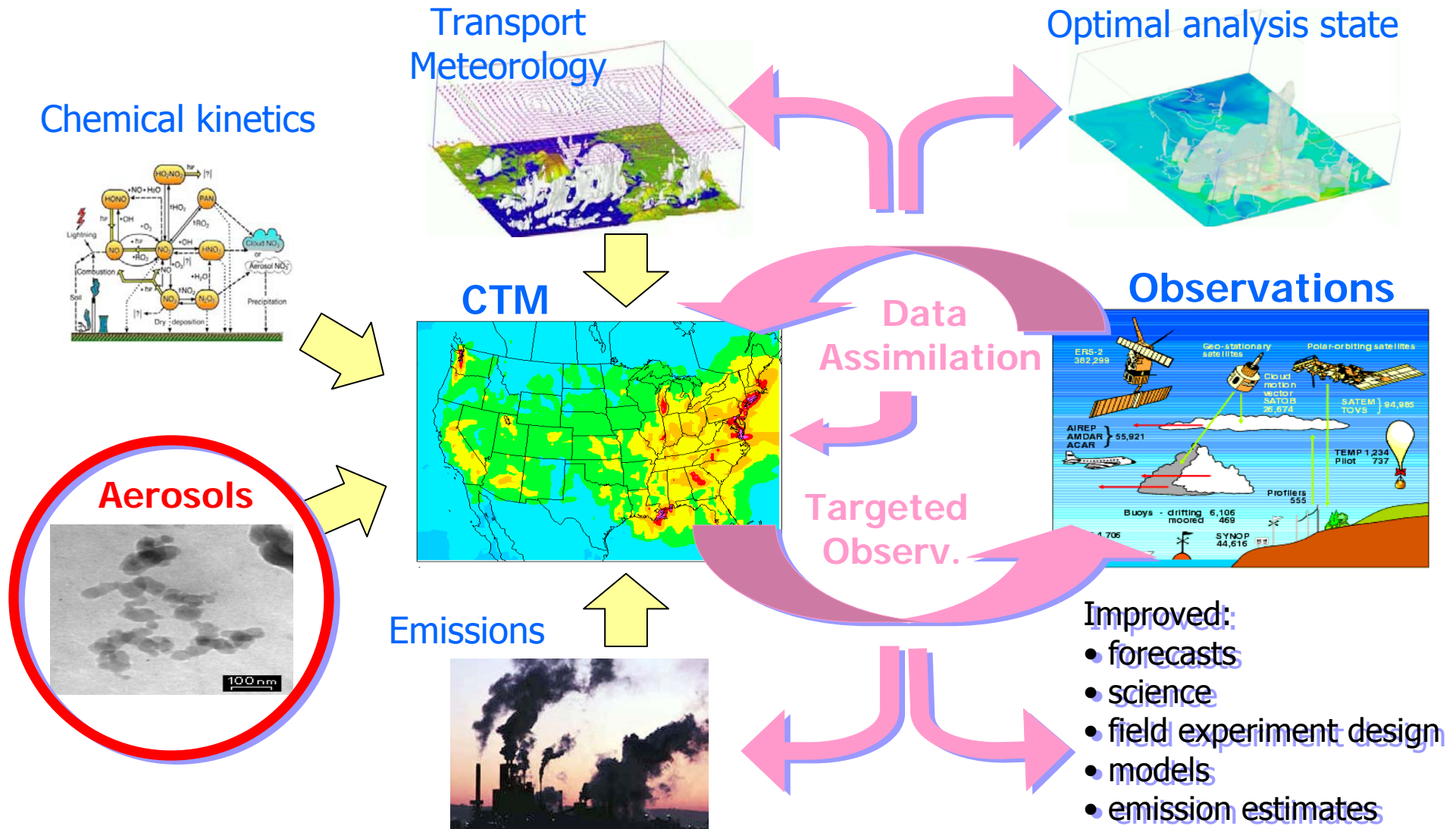
## TLM (DDM)



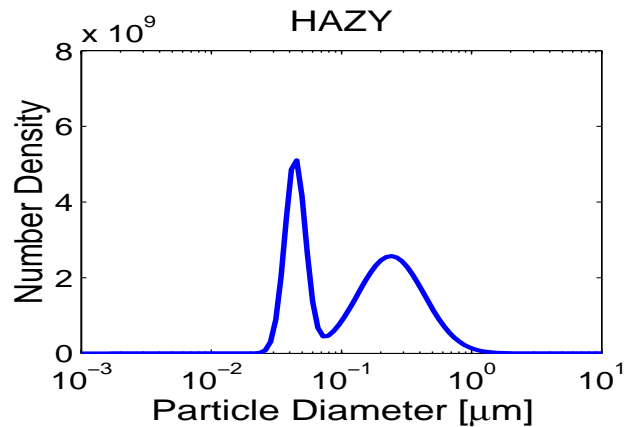
## Discrete ADJ



# Adjoins for Integral-PDE aerosol dynamic equations: formulation and challenges



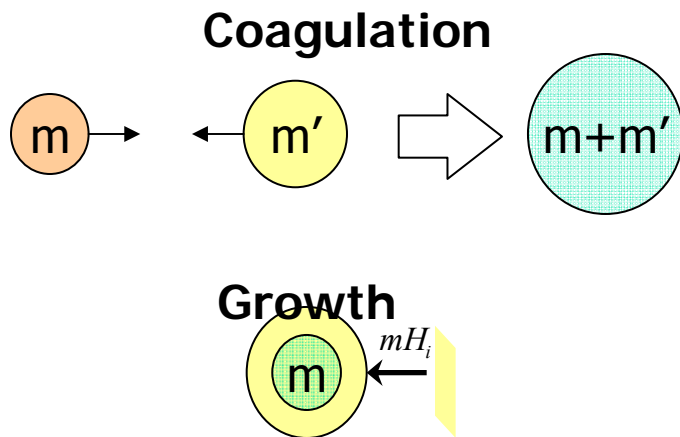
# Populations of aerosols (particles in the atmosphere) are described by their mass density



## Aerosol dynamic equation - IPDE

$$\frac{\partial q_i}{\partial t} = \int_0^m \beta(m', m - m') q_i(m', t) \frac{q(m - m', t)}{m - m'} dm' - q_i \int_0^\infty \beta(m, m') \frac{q(m', t)}{m'} dm' + H_i q - \frac{\partial}{\partial m} (m H_i q_i) + m_i S - L q_i + R_i(q)$$

$$q_i(m, t = t^0) = q_i^0(m), \quad 1 \leq i \leq n, \\ q_i(m = 0, t) = 0, \quad q_i(m = \infty, t) = 0.$$



# Adjoint aerosol dynamic models are needed to solve inverse problems

$$\frac{\partial \lambda_i}{\partial t} = - \int_0^{\infty} \beta(m, m') (m')^{-1} [\lambda_i(m + m', t) - \lambda_i(m, t)] q(m', t) dm' + L \lambda_i \quad t_+^{k-1} \leq t \leq t_-^k$$

Continuous  
adjoint  
equation

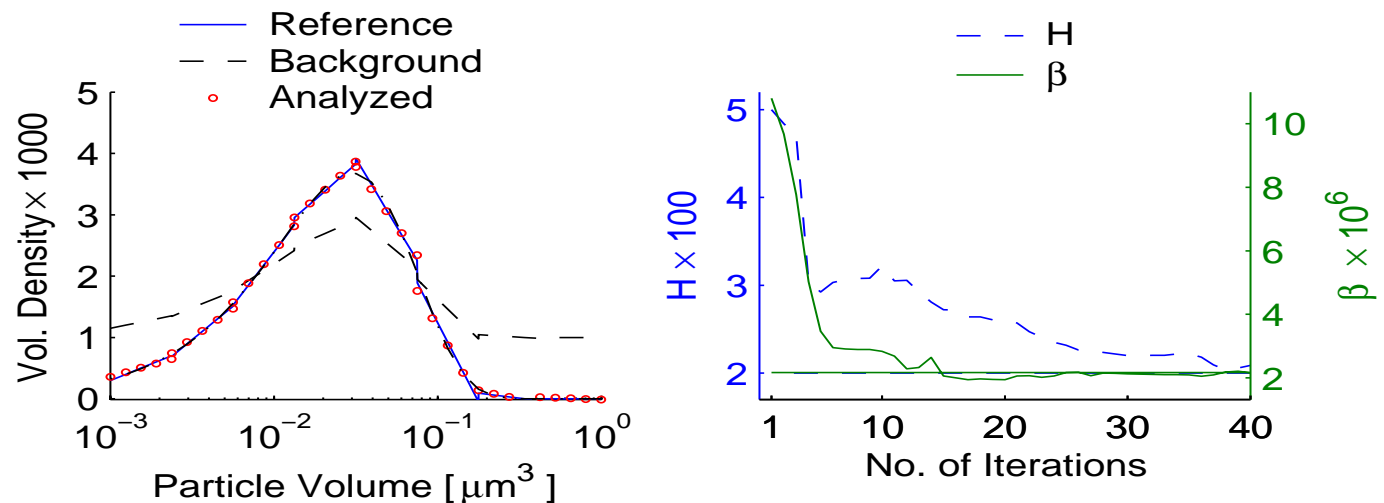
$$- \int_0^{\infty} \beta(m', m) m^{-1} \sum_{j=1}^n [\lambda_j(m + m', t) - \lambda_j(m, t)] q_j(m', t) dm' - \sum_{j=1}^n H_j \lambda_j - m H \frac{\partial \lambda_i}{\partial m}$$

$$\lambda_i(m, t^N) = 0, \quad \lambda_i(m, t_-^k) = \lambda_i(m, t_+^k) + h_{q_i}^T R_k^{-1} (y^k - h(q^k))$$

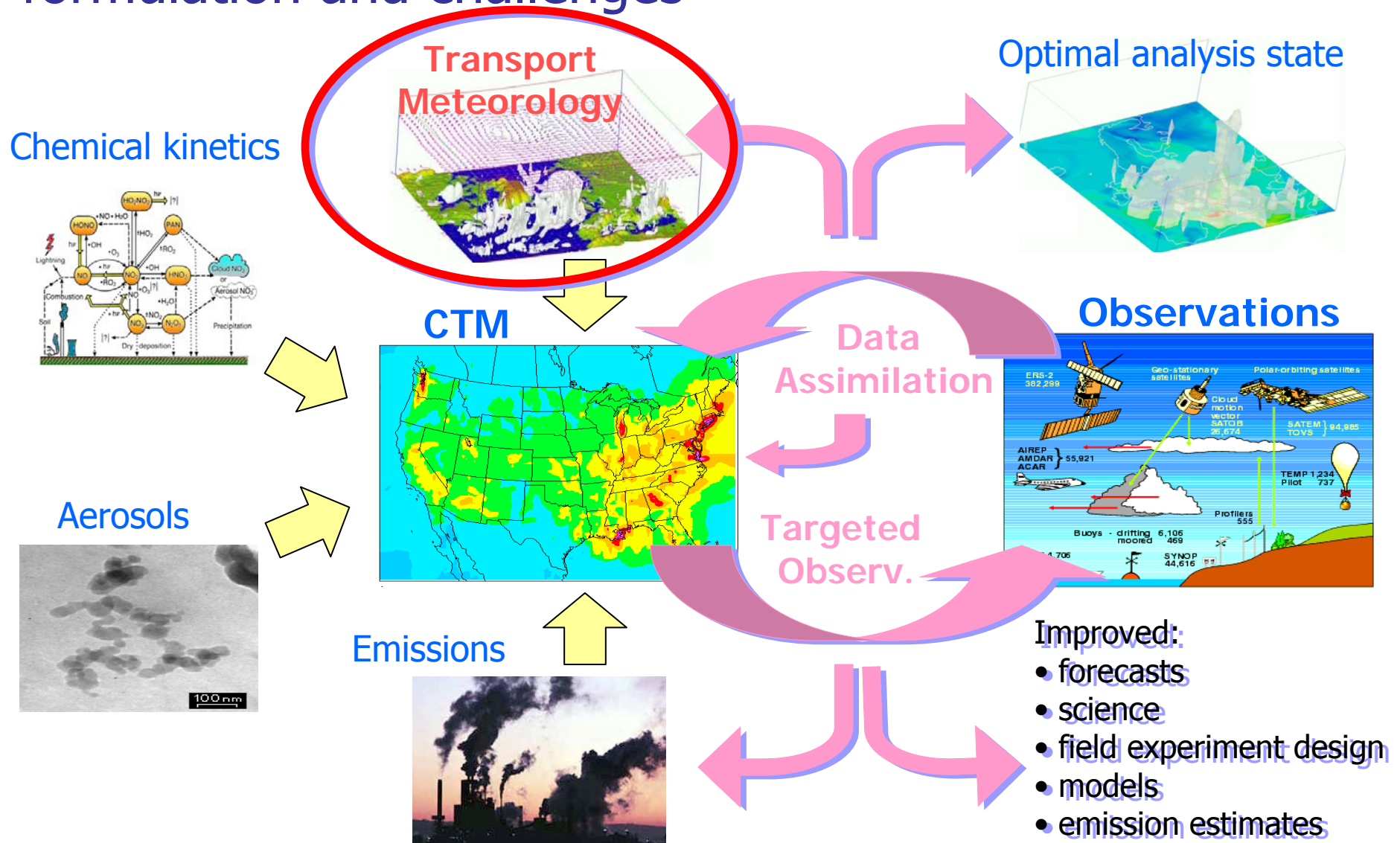
$$\lambda_i(m, t^0) = \lambda_i(m, t_+^0) + p_{q_i}^T B^{-1} (p - p^B), \quad \lambda_i(0, t) = 0.$$

Observations of density in each bin allow the recovery of initial distribution and of parameters

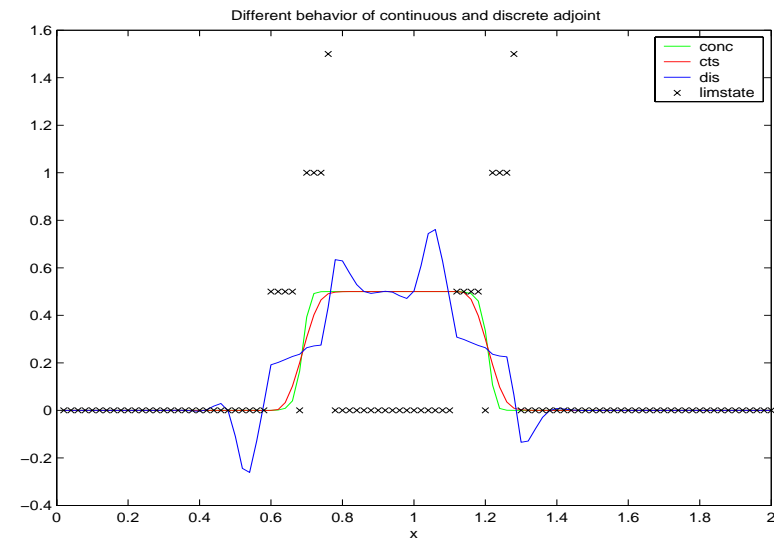
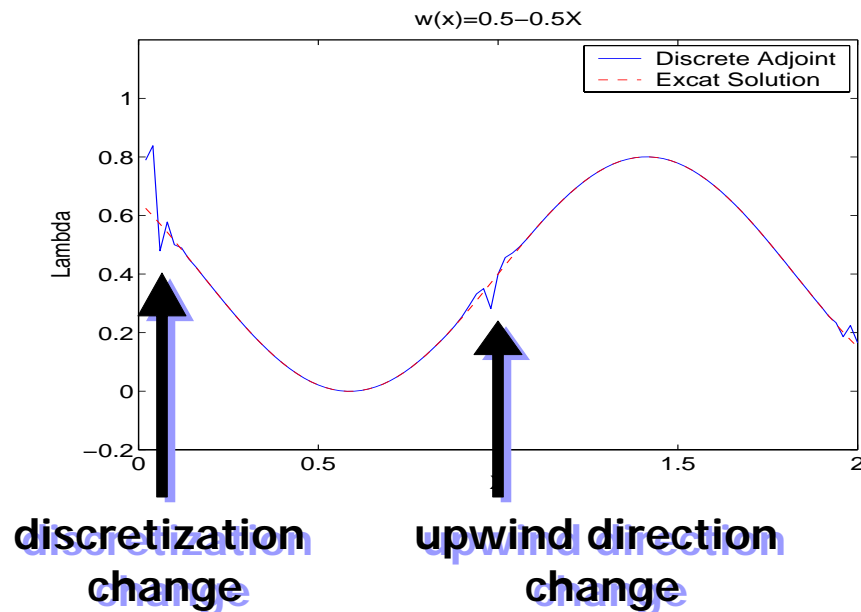
[Sandu et al., 2005;  
Henze et al., 2004]



# Discrete adjoint models for numerical advection: formulation and challenges



# Discrete adjoints of advection numerical schemes can become pointwise inconsistent with the adjoint PDE



Change of forward scheme pattern:

- Change of upwinding
- Sources/sinks
- Inflow boundaries scheme

Example: 3<sup>rd</sup> order upwind FD

Active forward limiters  
act as pseudo-sources in adjoint  
Example: minmod

*[Liu and Sandu, 2005]*

# The 4D-Var tools have been implemented in parallel adjoint STEM and are being applied to real data

## Chemistry: KPP

- Forward: sparse Rosenbrock, RK, LMM
- DDM sensitivity (Rosenbrock, LMM)
- Discrete adjoints: Rosenbrock
- Continuous adjoints: Rosenbrock, RK, LMM

## Aerosols: 0-D, not yet 3-D

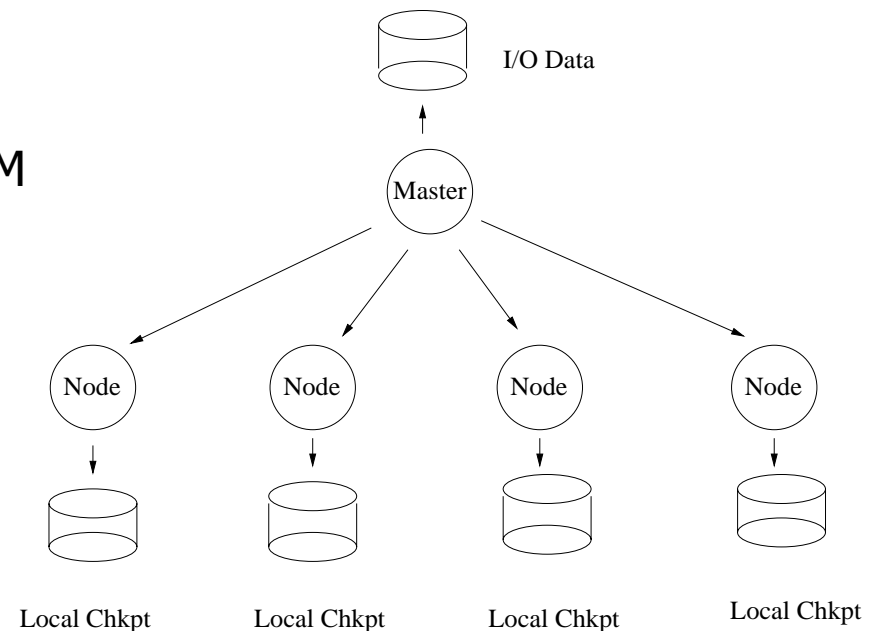
## Transport:

- Forward: upwind FV, FD, FE
- Adjoint for linear upwind FD

## Parallelization: with PAQMSG

*[Sandu et.al., 2003, 2004; Carmichael et. al., 2003, 2004]*

## Distributed, 2-level checkpointing scheme



# Dynamic integration of chemical data and atmospheric models is an important, growing field

## ■ **Current state of the art:**

- the tools needed for 4d-Var chemical data assimilation are in place:  
(adjoints for stiff systems, aerosols, transport; singular vectors, parallelization and multi-level checkpointing schemes, models of background errors)
- their strengths demonstrated using real (field campaign) data; ambitious science projects are ongoing
- ensemble-based chemical data assimilation is new, but promising

## ■ **Computational tools are being widely adopted:**

EPA (CMAQ), JPL (GEOS-Chem), NCAR and NOAA (WRF-Chem), Canadian Meteorological Centre, Max Planck Institute Germany (MESSY), University of Koeln (EURAD - EUROpean Air Pollution Dispersion)